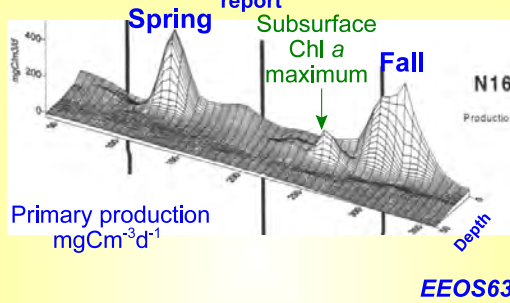


<div> <div> Light effects on photosynthesis, Model P vs. I approach for estimating production & Subsurface Chlorophyll maxima </div> <div> Class 18, 10/30/08 </div> <div> EEOS630 </div> </div>	<div>Slide 1 Light effects on photosynthesis, Model P vs. I approach for estimating production & Subsurface Chlorophyll maxima</div> <div>NOTES:</div>
<div> <div>Wimba Sessions</div> <div> <p>Tonight, Thursday 10/30, 7 pm – only if I can get files posted - Will know by 3 pm.</p> <ul style="list-style-type: none"> Quantitative community analysis using Matlab <ul style="list-style-type: none"> I've had to do rewriting of my Matlab m.files. Run the tutorial at the Mathworks site I'll be logged on at 7 pm tonight to demonstrate community analysis using the West Falmouth oilspill data as an example using Matlab I'll also present the solution to the <i>Synechococcus</i> problem Due date: papers due 4 weeks after projects posted. </div> <div>EEOS630</div> </div>	<div>Slide 2 Wimba Sessions</div> <div>NOTES:</div>
<div> <div>Phytoplankton Readings</div> <div> <div>Nutrients and the spring bloom</div> <ul style="list-style-type: none"> Nutrient effects, Tuesday <ul style="list-style-type: none"> Chapter 10: Nitrogen cycle, nutrient limitation & chemostats Howarth, R. W. 1988. Nutrient limitation of net primary production in marine ecosystems. <i>Ann. Rev. Ecol. Syst.</i> 19: 89-110. Spring bloom, Thursday <ul style="list-style-type: none"> Chapter 11: Sverdrup's critical depth concept & the vernal phytoplankton Sverdrup, H. U. 1953. On conditions for the vernal blooming of phytoplankton. <i>J. Conseil perm. int. Explor. Mer.</i> 18: 287-295. Parsons, T. R., M. Takahashi, and B. Hargrave. 1984. <i>Biological Oceanographic Processes</i>. 3rd Edition. Pergamon Press, Oxford & New York. Pages 87-100. Townsend, D. W. and R. W. Spinrad. 1986. Early phytoplankton blooms in the Gulf of Maine. <i>Cont. Shelf Res.</i> 6: 515-529. </div> </div>	<div>Slide 3 Phytoplankton Readings</div> <div>NOTES:</div>

1995 MA Bay seasonal production

Craig Taylor's data in MWRA 1995 water-column report

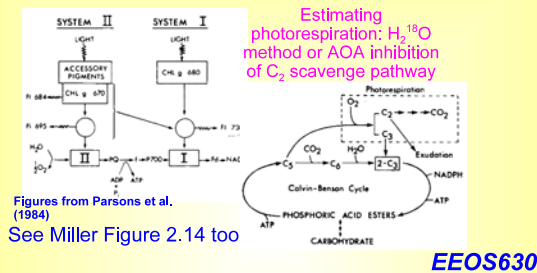


Slide 4 1995 MA Bay seasonal production

NOTES:

Gross primary Productivity

Light-dependent rate of electron flow to terminal electron acceptors (Falkowski & Raven, 1997 p 264)

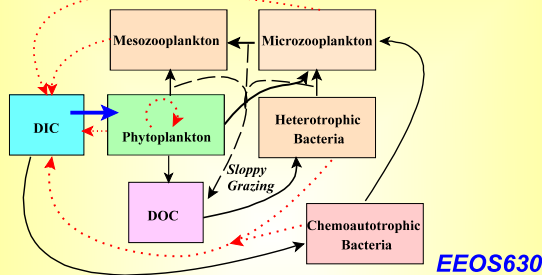


Slide 5 Gross primary Productivity

NOTES:

Carbon flow in bottles (& the sea)

Are short incubations the best solution?
Photoacclimation a problem



Slide 6 Carbon flow in bottles (& the sea)

NOTES:

The oxygen method

Separating Gross vs. Net production.

- Use light and dark bottles
- Dark bottle measures respiration (but not photorespiration)
 - Photorespiration may not be too important
 - In the absence of grazers, light bottle measures net production.
 - In the absence of grazers, light-Dark measures gross production.
- Photorespiration and heterotrophs create large problems

EEOS630

Slide 7 The oxygen method

NOTES:

O₂ vs. ¹⁴C

$n \text{ CO}_2 \xrightarrow{\text{light}} \underset{1}{\text{Particulate organic C}} + \underset{2}{\text{Dissolved organic C}} + \underset{3}{n \text{ O}_2}$

O₂ method measures O₂ production, = measuring 1 + 2.
¹⁴C method measures only 1 if only filtered POC is counted.

- In theory, the O₂ method can estimate gross and net production
 - Increase in light bottle is net
 - Dark bottle decrease is respiration
 - Light-dark = gross production
- The ¹⁴C activity of dissolved organic matter should be determined

EEOS630

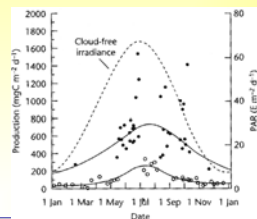
Slide 8 O₂ vs. ¹⁴C

NOTES:

Primary production underestimated

Miller (2004) Fig. 3.4, Welschmeyer *et al.* 1993

- Open circles: data collected before 1980
- Filled circles 1980-1984, Trace-metal clean conditions
- 2-fold difference
- There may have also been interannual variability: the Pacific interdecadal oscillation (first discovered after Welschmeyer's 1993 paper)



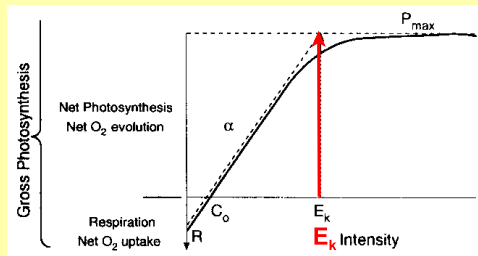
ECOS630

Slide 9 Primary production underestimated

NOTES:

Falkowski & Raven P vs. E curves

Falkowski & Raven (1997, p. 196, Fig 7.2)

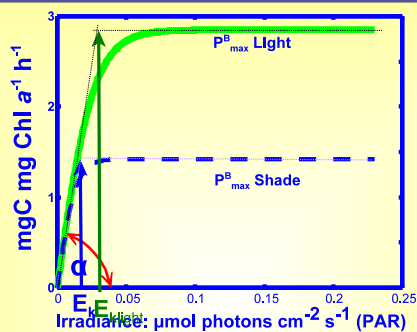


EEOS630

Slide 10 Falkowski & Raven P vs. E curves

NOTES:

Chl a-specific gross productivity



Slide 11

NOTES:

PAR & units of light intensity

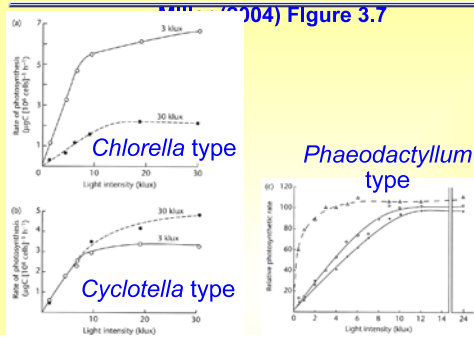
- Parsons, Takahashi & Hargrave (1984)
 - Parsons_1984.pdf on ereserve
 - the direct link the ECOS630 reserves is: <http://docutek.lib.umb.edu/eres/coursepass.aspx?cid=65>
 - Password: deep
- Photosynthetically available radiation [PAR] The quantity of light in those wavelengths that can be utilized for photosynthesis (400 to 700 nm Behrenfeld & Falkowski (1997), some earlier papers listed PAR from 300 to 720 nm);
 - Falkowski & Raven (2007, 341), PAR was defined as 350 to 720, but it is difficult to construct detectors to measure photon flux at 340-400 nm, so now it is 400 to 700)
- To convert to energy, 550nm light assumed

EEOS630

Slide 12 PAR & units of light intensity

NOTES:

Types of shade acclimation



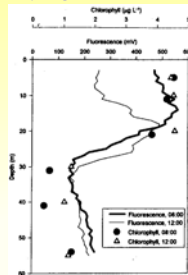
Slide 16 Types of shade acclimation

NOTES:

Measuring the Chl a profile

Falkowski & Raven (1997) Figure 9.6

- **Fluorescence yield reduced by:**
 - Photochemical quenching: can occur with a diel cycle
 - Shade acclimation & increase in PSII activity
 - Increase in the physiological status of the cells
- **Fluorescence yield Increased by:**
 - Reduction in PSII activity
 - Toxic effects (used to estimate Cu contamination)
 - Nutrient stress
 - Senescence



EEOS630

Slide 17 Measuring the Chl a profile

NOTES:

Shade acclimation (adaptation)

Phytoplankton can adapt to both the intensity and quality of light by:

- Changes in the ratios of photosynthetic pigments
- Changes in the size and number of photosynthetic units (measured by the Chl a:P700 size, Perry *et al.* 1981)
- Changes in chloroplast size & orientation
- Changes in the enzyme activities of both the light and dark reactions

EEOS630

Slide 18 Shade acclimation (adaptation)

NOTES:

Quenching

Refers to both scintillation counting and fluorescence yield; Falkowski & Raven (2007)

Quenching and quantum yield for fluorescence

$$\Phi_F = \frac{k_F}{(k_F + k_A + k_P)}, \text{ where}$$

Φ_F = ratio of light emitted as fluorescence to light absorbed.

k_F = de-excitation of singlet to fluorescence

k_A = heat.

k_P = photochemistry.

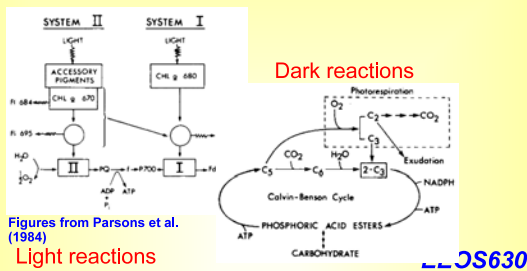
EEOS630

Slide 19 Quenching

NOTES:

Gross primary productivity

Light-dependent rate of electron flow to terminal electron acceptors (Falkowski & Raven, 1997 p 264)



Slide 20 Gross primary productivity

NOTES:

Jassby & Platt's (1976) Equation

Without photoinhibition

$$P^II = P_{II}^{max} \tanh\left(\frac{\alpha I}{P_{II}^{max}}\right)$$

where P^II = $C H_2 a$ -specific production $\left[\frac{mg C}{mg Chla h} \right]$

P_{II}^{max} = Max. rate at light saturation.

= Assimilation number

= the maximum photosynthetic rate at light saturation.

α = initial slope of the P vs. I curve.

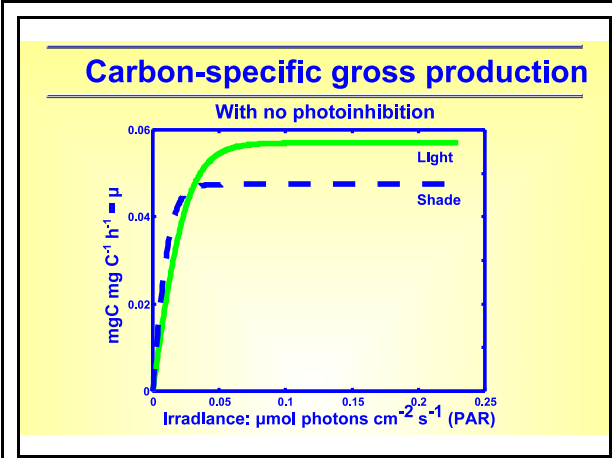
= $\left[\frac{mg C}{mg Chla h (W m^{-2})} \right]$ or

= $\left[\frac{mg C}{mg Chla h (\mu E m^{-2} s^{-1})} \right]$

I = the light intensity of PAR $[W m^{-2}]$ or $[Einstein m^{-2} s^{-1}]$.

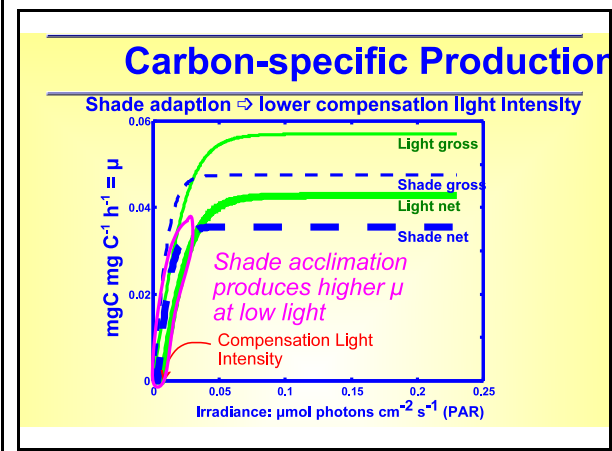
Slide 21 Jassby & Platt's (1976) Equation

NOTES:



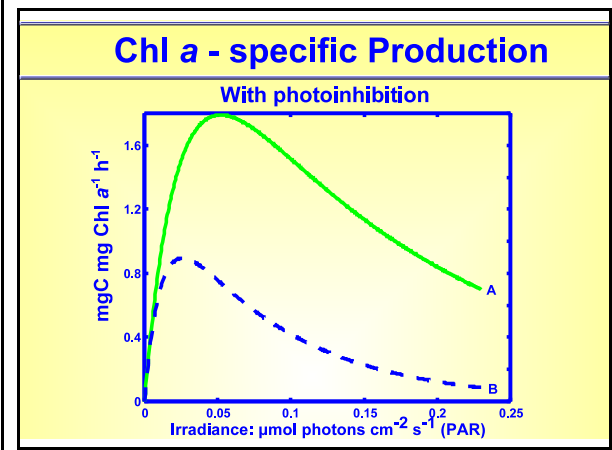
Slide 22 Carbon-specific gross production

NOTES:



Slide 23 Carbon-specific Production

NOTES:



Slide 24 Chl a - specific Production

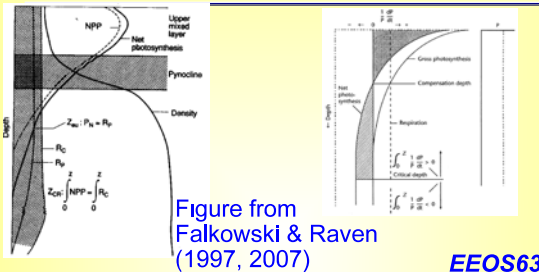
NOTES:

<div data-bbox="334 170 704 205" data-label="Section-Header"><p>Photoinhibition Equation</p></div> <div data-bbox="410 216 604 239" data-label="Text"><p>Jassby & Platt (1976)</p></div> <div data-bbox="256 275 764 451" data-label="Equation-Block">$P^B = P_r^B \left(1 - e \left(\frac{-\alpha I}{P_r^B} \right) \right) e^{-\frac{\beta I}{P_r^B}}$<p>$\alpha$ - Initial slope of P vs. I curve. β - Photoinhibition parameter. - Negative slope at high light intensity. P_r^B - Max. photo. rate without photoinhibition.</p></div> <div data-bbox="662 516 769 541" data-label="Text"><p>EEOS630</p></div>	<div data-bbox="824 136 1416 178" data-label="Section-Header"><p>Slide 25 Photoinhibition Equation</p></div> <div data-bbox="824 258 938 289" data-label="Text"><p>NOTES:</p></div>
<div data-bbox="272 659 773 695" data-label="Section-Header"><p>Carbon-specific gross production</p></div> <div data-bbox="297 705 688 1031" data-label="Figure"><p>With photoinhibition</p><p>The graph shows carbon-specific gross production (mgC mg C⁻¹ h⁻¹ = μ) on the y-axis (0 to 0.04) versus irradiance (μmol photons cm⁻² s⁻¹ (PAR)) on the x-axis (0 to 0.25). A solid green curve labeled 'Light' rises to a peak of approximately 0.035 at an irradiance of about 0.05, then declines. A dashed blue curve labeled 'Shade' follows the initial rise of the 'Light' curve and then levels off at a lower value of about 0.015 at higher irradiances.</p></div>	<div data-bbox="824 625 1416 667" data-label="Section-Header"><p>Slide 26 Carbon-specific gross production</p></div> <div data-bbox="824 747 938 779" data-label="Text"><p>NOTES:</p></div>
<div data-bbox="272 1148 769 1184" data-label="Section-Header"><p>Vertical profiles of photosynthesis</p></div> <div data-bbox="248 1194 721 1486" data-label="Figure"><p>Miller (2004) Fig. 3.9, May (●) & Sept (○)</p><p>Two side-by-side graphs showing vertical profiles from 0 to -90 meters depth. Graph (a) shows 'Productivity (mgC m⁻² d⁻¹)' on the x-axis (0 to 40). Graph (b) shows 'Chlorophyll a (μg L⁻¹)' on the x-axis (0 to 0.6). Both graphs show data for May (closed circles) and September (open circles). In both, the highest values are near the surface (0-10m) and then decrease with depth. Productivity peaks are around 30-40 mgC m⁻² d⁻¹ in May and 10-20 in September. Chlorophyll a peaks are around 0.4-0.5 μg L⁻¹ in May and 0.2-0.3 in September.</p></div> <div data-bbox="248 1488 721 1526" data-label="Caption"><p>Fig. 3.9 (a) May (closed circles) and September (open circles) profiles of primary productivity m⁻² in the Gulf of Alaska (59°N, 143°W). (b) Chlorophyll profiles from the same site at various times in the same months (same symbols are always maximal near the surface, then taper off below 50 m. (After Wischniwer 1993.)</p></div>	<div data-bbox="824 1115 1416 1157" data-label="Section-Header"><p>Slide 27 Vertical profiles of photosynthesis</p></div> <div data-bbox="824 1236 938 1268" data-label="Text"><p>NOTES:</p></div>



<div data-bbox="207 132 792 573"> <h3>Hourly Gross Productivity vs. Depth</h3> <p>Light Intensity: $I_z = I_o \cdot \exp(-K_d \cdot z)$; Beer's law $\xi = K_d \cdot z = \text{optical depth}$</p> <p>A) Early morning B) Mid Morning C) Noon</p> </div>	<div data-bbox="824 132 1414 216"> <h3>Slide 28 Hourly Gross Productivity vs. Depth</h3> </div> <div data-bbox="824 289 1414 342"> <p>NOTES:</p> </div>
<div data-bbox="207 663 792 1104"> <h3>Hourly Gross Productivity</h3> <p>With photoinhibition</p> <p>PAR Surface Beneath surface Lower photic zone</p> </div>	<div data-bbox="824 663 1414 705"> <h3>Slide 29 Hourly Gross Productivity</h3> </div> <div data-bbox="824 779 1414 831"> <p>NOTES:</p> </div>
<div data-bbox="207 1146 792 1587"> <h3>Cells are not stationary!</h3> <p>Left: position of cell; right: light history of cells at 2 depths</p> <p>No mixing High mixing</p> <p>EEOS630</p> </div>	<div data-bbox="824 1146 1414 1188"> <h3>Slide 30 Cells are not stationary!</h3> </div> <div data-bbox="824 1262 1414 1314"> <p>NOTES:</p> </div>

Euphotic zone ($\approx 1\%$ light depth), mixed layer depth, and critical depth



Slide 31 Euphotic zone ($\approx 1\%$ light depth), mixed layer depth, and critical depth

NOTES:

Non-dimensional production

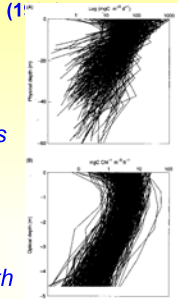
Behrenfeld & Falkowski (1997), Falkowski & Raven (1997)

Production vs. Depth

Meters

Chl-specific P vs. Depth

Optical depth



Slide 32 Non-dimensional production

NOTES:

The model P vs I approach vs. SIS

Harrison et al. (1985)

- Light & Chl a profiles determined
- SIS
 - ▶ 30-L Niskin bottles from 100, 50, 25, 10, and 1% light depths
 - ▶ 24 h 200-ml incubations on board deck
 - ▶ natural sunlight & neutral density filters
 - ▶ Trapezoidal integration over depth interval
- Model P vs. I approach
 - ▶ Water collected at 50% and 1% Light depths
 - ▶ 30-50 100-ml incubations
 - ▶ Artificial light
 - ▶ Time-zero or dark-bottle blanks

EEOS630

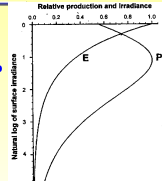
Slide 33 The model P vs I approach vs. SIS

NOTES:

Model estimates of production

Requires profiles of light and Chl *a*

- Profile light over a 24-h day and chl *a*
- Estimate α and AN at the 1% and 50% light depth only using 30-50 P vs. I incubations
- Calculate hourly production using Chl *a*, I, P_0 and α
- Sum 24-h estimates to obtain daily production
- Compare with SIS profiles



$$P(t) = P_0 (I(t))^\alpha \cdot E$$

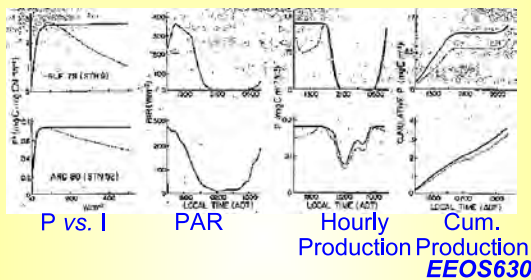
$$P_d = \sum_{i=1}^{24} P(t_i)$$

Slide 34 Model estimates of production

NOTES:

Model P vs. I approach

Fig 1 (dotted=photoinhibition) Harrison *et al.* (1985)

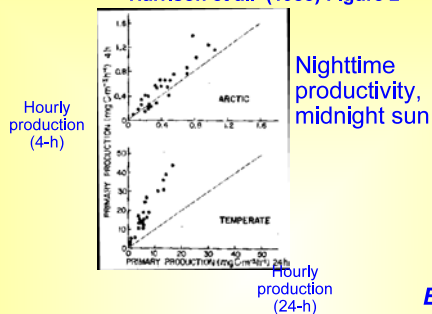


Slide 35 Model P vs. I approach

NOTES:

4- vs. 24-h incubations

Harrison *et al.* (1985) Figure 2



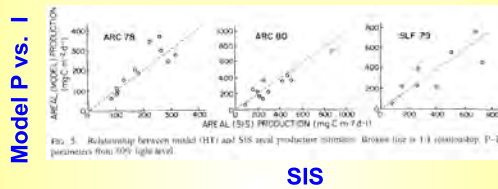
Slide 36 4- vs. 24-h incubations

NOTES:

<p style="text-align: center;">Harrison et al. (1985)</p> <p style="text-align: center;">Fig. 3. P vs. I parameters from 50% Light Depth</p> <p style="text-align: center;">Solid line: SIS Broken: No β Dotted: With β</p> <p style="text-align: right;">EEOS630</p>	<p>Slide 37 Harrison et al. (1985)</p> <p>NOTES:</p>
<p style="writing-mode: vertical-rl; transform: rotate(180deg);">Model P vs. I approach</p> <p style="text-align: center;">SIS approach</p> <p style="text-align: right;">EEOS630</p>	<p>Slide 38</p> <p>NOTES:</p>
<p style="text-align: center;">Problems at low light intensities</p> <p style="text-align: center;">The model underestimates SIS production.</p> <p style="writing-mode: vertical-rl; transform: rotate(180deg);">Model P vs. I</p> <p style="text-align: center;">SIS More blue light</p> <p style="text-align: right;">EEOS630</p>	<p>Slide 39 Problems at low light intensities</p> <p>NOTES:</p>

Areal production accurate

Harrison et al. (1985) Fig. 5



EEOS630

Slide 40 Areal production accurate

NOTES:

P vs. I parameters indicate shade adaptation/stratification

50% Light depth chosen for P vs. I parameters

TABLE 3. Depth differences in P-I model parameters (units given in text). *Depth difference significant at the 95% confidence level, **significant at the 99% confidence level.

Cruise	No. of observations	Light depth (%)	P-I parameters: $f \pm \text{SE}$			
			P_s	P_m	α	β
ARC 78	12	50	1.50 (0.16)	1.42 (0.15)**	0.052 (0.003)**	0.0005 (0.0001)**
	12	1	1.34 (0.18)	0.84 (0.12)	0.039 (0.005)	0.0056 (0.0010)
ARC 80	8	50	1.20 (0.28)	1.07 (0.24)*	0.054 (0.010)*	0.0012 (0.0003)
	8	1	0.88 (0.11)	0.76 (0.10)	0.047 (0.008)	0.0046 (0.0004)
SLF 79	8	50	3.46 (0.49)	2.80 (0.31)**	0.056 (0.008)	0.0035 (0.0014)*
	8	1	3.25 (0.56)	1.91 (0.34)	0.044 (0.007)	0.0063 (0.0023)

Shade acclimation

EEOS630

Slide 41 P vs. I parameters indicate shade adaptation/stratification

NOTES:

Applications

- Application to MA Bay monitoring
 - Measuring Chl a
 - Measuring production
 - O_2
 - Model approach introduced by Craig Taylor
 - Incubation method
 - Problems with the model approach
 - Assimilation numbers too high
 - Subsurface chlorophyll maximum can be a productivity maximum in MA Bay
 - Which P vs. I parameters should be used?

EEOS630

Slide 42 Applications

NOTES:

The MA Bay Outfall

From MWRA State of the Harbor Report



Figure 9a Map of Massachusetts - Boston Harbor



Figure 9b Map of Massachusetts - Cape Cod Bay

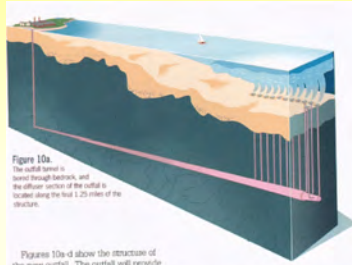


Figure 10a-d show the structure of the new outfall. The outfall will provide

EEOS630

Slide 43 The MA Bay Outfall

NOTES:

Realistic P vs. I parameters

Many published assimilation numbers and α 's are too high

- There **are** theoretical limits for α and A.N.
 - ▶ Maximum α set by the quantum efficiency of photosynthesis
 - ▶ A.N. is set by the maximum specific growth rate (assuming balanced growth)
 - ▶ Falkowski published a theoretical maximum of about 20-25
 - ▶ Harris in Phytoplankton Ecology, maximum assimilation numbers in the gyres should be: 6-8 mg C mg Chl $a^{-1} h^{-1}$

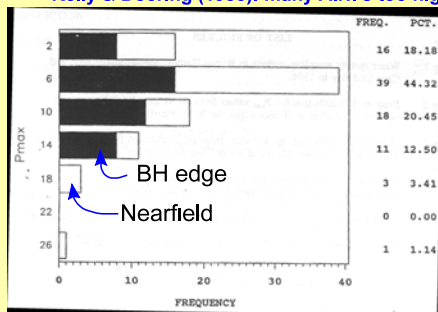
EEOS630

Slide 44 Realistic P vs. I parameters

NOTES:

MA Bay P vs. I parameters

Kelly & Doering (1985): many A.N.'s too high!



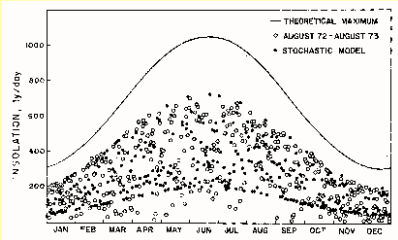
OS630

Slide 45 MA Bay P vs. I parameters

NOTES:

Model P vs. I approach: can account for differences in irradiance (due to clouds)

Rhode Island solar irradiance (from Kremer & Nixon)



EEOS630

Slide 49 Model P vs. I approach: can account for differences in irradiance (due to clouds)

NOTES:

1995 MA Bay Production

Craig Taylor (WHOI) Model P vs. I approach

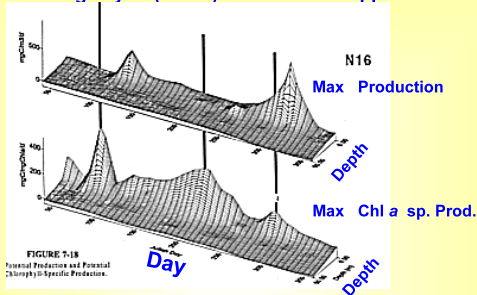


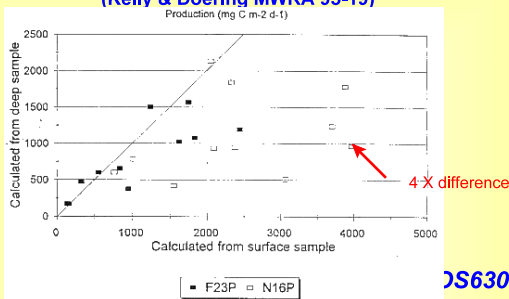
FIGURE 7-18
Potential Production and Potential
Chlorophyll-Specific Production.

Slide 50 1995 MA Bay Production

NOTES:

2 different production estimates

Using deep vs. surface P vs. I parameter estimates
(Kelly & Doering MWRA 95-19)



DS630

Slide 51 2 different production estimates

NOTES:

Excursis on the Subsurface Chlorophyll maxima

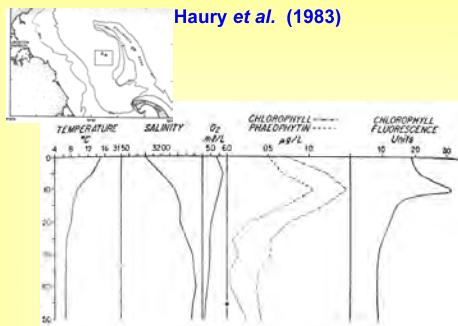
Prevalent in Gulf of Maine (and MA Bay) from April through early September, many coastal zones (including the Washington-Oregon-California shelf) & oligotrophic gyres (called the Typical Tropical Structure)

EEOS630

Slide 52 Excursis on the Subsurface Chlorophyll maxima

NOTES:

MA Bay subsurface Chl a maxima



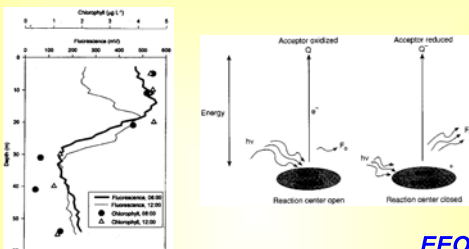
DS630

Slide 53 MA Bay subsurface Chl a maxima

NOTES:

As noted by Cullen, SSFluoresence not necessarily a SSChl max nor SSCarbon max

Falkowski & Raven Figure 9.6



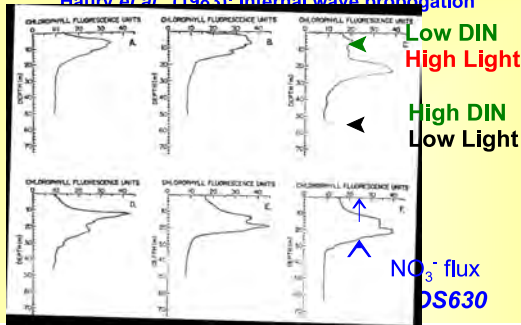
EEOS630

Slide 54 As noted by Cullen, SSFluoresence not necessarily a SSChl max nor SSCarbon max

NOTES:

Internal waves and MA Bay SSCM

Haury et al. (1983): internal wave propagation

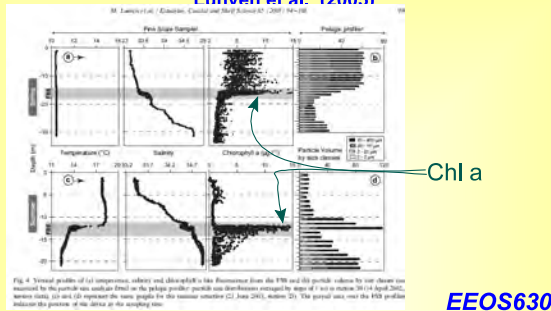


Slide 55 Internal waves and MA Bay SSCM

NOTES:

Fine structure of the SSCM

Lunven et al. (2005)

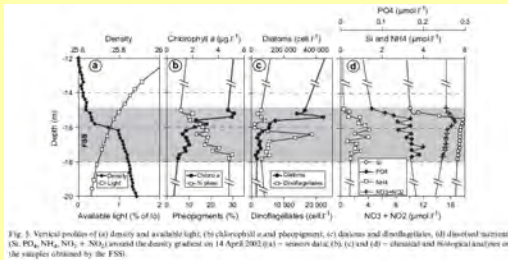


Slide 56 Fine structure of the SSCM

NOTES:

Fine structure of the SSCM

Lunven et al. (2005)

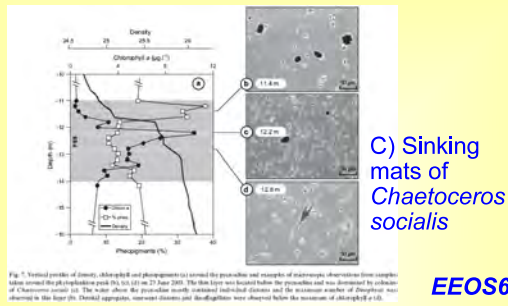


Slide 57 Fine structure of the SSCM

NOTES:

Fine structure of the SSCM

Lunven et al. (2005)

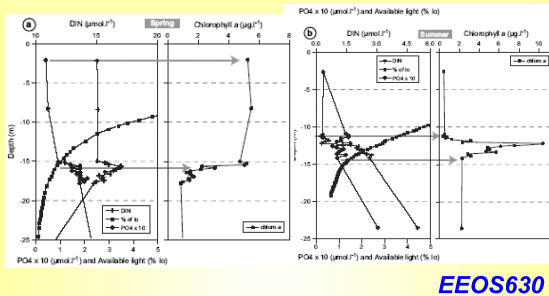


Slide 58 Fine structure of the SSCM

NOTES:

Fine structure of the SSCM

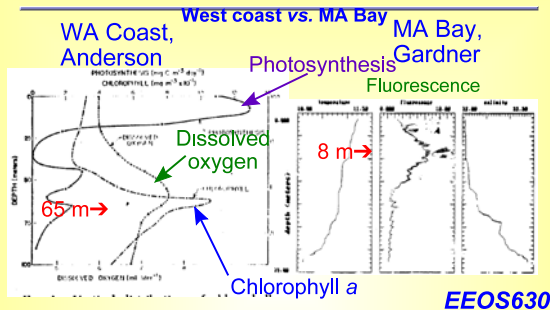
Lunven et al. (2005): 0.5% light level at base of SSCM



Slide 59 Fine structure of the SSCM

NOTES:

SSCM off the Washington-Oregon Coast, also off California

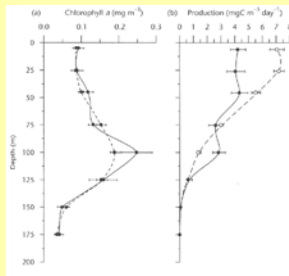


Slide 60 SSCM off the Washington-Oregon Coast, also off California

NOTES:

Central N. Pacific gyre: Typical tropical structure

SSCM at 100 meters; Miller (2004) Fig. 10.6



In Lundven's European coastal zone & in MA Bay, the SSCM can be a major component of total water-column production

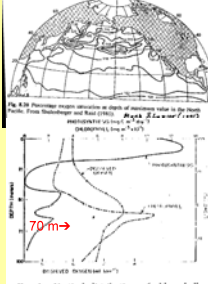
EEOS630

Slide 61 Central N. Pacific gyre: Typical tropical structure

NOTES:

SSCM O₂: 120% saturation

Shulenberger & Reid (1981), Jenkins (1982) in Atlantic



AOU: Apparent oxygen utilization at depth, convert to new production using Redfield ratio

EEOS630

Slide 62 SSCM O₂: 120% saturation

NOTES:

Nutrient limitation

EEOS630

Slide 63 Nutrient limitation

NOTES:

Four major revolutions

In our understanding of nutrient limitation

- Brandt (1899) was correct to focus on N limitation, Liebig's law, and the role of denitrification, but he missed the role of vertical mixing providing vertical flux of nutrients
 - The anammox pathway, missed until 2003 provides further insight into the central role of nitrogen removal
- Chemostat work by Droop (1968), Caperon & Meyer (1972), Fuhs & Rhee revealed the central importance of the **Internal nutrient pool** in controlling μ
- Goldman (Goldman *et al.* 1979, 1980) argued that phytoplankton in nature tend to grow at high relative growth rates, otherwise they would not exhibit Redfield stoichiometry. The internal nutrient pool tends to follow Redfield stoichiometry.
 - Nutrient input controls phytoplankton biomass & species composition
 - One phytoplankton assemblage rapidly replaced by another, each with high relative growth rate.
- Martin's Iron hypothesis: iron is the Liebigian nutrient in major areas of the world's ocean

EEOS630

Slide 64 Four major revolutions

NOTES:

Terms & concepts

Quick List

- Nitrogen cycle
 - nitrification
 - denitrification (dissimilatory nitrate reduction)
 - Assimilatory nitrate reduction
- Nutrient limitation: What are the different meanings
- Chemostats
 - What are they?
 - Michaelis-Menten Equation
 - Monod Equation
 - Droop Equation
 - Caperon & Meyer's (1972)'s equation
- Four major revolutions in understanding nutrient limitation

EEOS630

Slide 65 Terms & concepts

NOTES:

Hensen's Nets & major cruise

From Mills (1989): 50- μ m nets, 1889 National cruise



Slide 66 Hensen's Nets & major cruise

NOTES:

<div data-bbox="282 168 753 205" data-label="Section-Header"> <h3>Hensen's "Blood of the Ocean"</h3> </div> <div data-bbox="339 214 693 241" data-label="Section-Header"> <h4>Uniformly distributed phytoplankton!</h4> </div> <div data-bbox="238 243 748 522" data-label="List-Group"> <ul style="list-style-type: none"> • The German Hensen introduced quantitative Plankton sampling to oceanography (1840s-1880s) <ul style="list-style-type: none"> ▸ Hensen introduced quantitative plankton sampling (50-μm silk mesh) ▸ Phytoplankton are uniformly distributed • Conclusions from 1889 National Cruise <ul style="list-style-type: none"> ▸ Within a biogeographic province, phytoplankton are uniformly distributed in the ocean, like oxygen and other chemical constituents ▸ The oceans were in general very poor in plankton standing stocks, especially the tropics. </div> <div data-bbox="656 514 771 541" data-label="Text"> <p>EEOS630</p> </div>	<div data-bbox="816 134 1375 170" data-label="Section-Header"> <h3>Slide 67 Hensen's "Blood of the Ocean"</h3> </div> <div data-bbox="816 258 940 291" data-label="Text"> <p>NOTES:</p> </div>
<div data-bbox="300 655 716 699" data-label="Section-Header"> <h3>SeaWiFS Average Chl a</h3> </div> <div data-bbox="397 705 609 732" data-label="Text"> <p>Oct 1997 -April 2002</p> </div> <div data-bbox="230 726 758 751" data-label="Text"> <p>http://seawifs.gsfc.nasa.gov/SEAWIFS/IMAGES/SEAWIFS_GALLERY.html</p> </div> <div data-bbox="224 753 764 1016" data-label="Figure"> </div> <div data-bbox="656 1005 771 1031" data-label="Text"> <p>EEOS630</p> </div>	<div data-bbox="816 621 1289 659" data-label="Section-Header"> <h3>Slide 68 SeaWiFS Average Chl a</h3> </div> <div data-bbox="816 743 940 777" data-label="Text"> <p>NOTES:</p> </div>
<div data-bbox="261 1144 786 1182" data-label="Section-Header"> <h3>Brandt's denitrification hypothesis</h3> </div> <div data-bbox="417 1188 589 1215" data-label="Text"> <p>Proposed in 1899</p> </div> <div data-bbox="238 1218 758 1497" data-label="List-Group"> <ul style="list-style-type: none"> • Nitrogen is the Liebigian (1876) limiting nutrient in the ocean <ul style="list-style-type: none"> ▸ Liebig proposed 50 agricultural laws, the law of the minimum was #33 (de Baar 1994) ▸ The law was proposed for monospecific crops • Why (according to Brandt, 1899)? <ul style="list-style-type: none"> ▸ Land is the major source of nitrogen to the sea ▸ Denitrifying bacteria have higher activities at higher temperatures ▸ Nitrogen should be scarcer in warmer waters ▸ Phytoplankton production should be less in tropical waters. </div> <div data-bbox="656 1488 771 1518" data-label="Text"> <p>EEOS630</p> </div>	<div data-bbox="816 1110 1276 1184" data-label="Section-Header"> <h3>Slide 69 Brandt's denitrification hypothesis</h3> </div> <div data-bbox="816 1268 940 1304" data-label="Text"> <p>NOTES:</p> </div>

The refutation of Brandt

Mixing & methodological problems

- Brandt assumed a terrestrial source for N.
Terrestrial sources are not the major source of Nitrogen fueling coastal production
 - the Norwegians proposed vertical mixing from deep, N-rich water
 - More vertical mixing in coastal waters, less in the tropics
- Methodological problems:
 - DIN could not be measured (until the 20s & 30s)
 - Denitrifying activity not measured until the 70s (Seitzinger)
 - Nitrifying bacterial activity not measured accurately until the 80's (Olson, Ward)

EEOS630

Slide 70 The refutation of Brandt

NOTES:

Brandt abandons his hypothesis

In 1929, opts for vertical mixing (Mills 1989, p. 161)

"The explanation is so evident that my explanation of 1899 that denitrifying bacteria are the cause of plankton deficiency in the tropical oceans is invalidated by it. However, I still maintain the view **"that denitrifying bacteria break down an excess of nitrogen compounds and that it is they that maintain the existing equilibrium in nature."**

EEOS630

Slide 71 Brandt abandons his hypothesis

NOTES:

Brandt's strengths & weaknesses

- Brandt was correct, but before his time in emphasizing:
 - Liebig's law of the minimum.
 - This has been tested experimentally, and it is usually only 1 nutrient, a rate-limiting nutrient that controls primary production
 - Multiple nutrient limitation not a major factor
 - Denitrification
 - Largely responsible for low N:P in marine waters
 - Phosphorus may be a limiting nutrient over geologic time scales & during glacial periods (Fe & N fixation)
- Major flaws
 - Overestimated terrestrial input of nitrogen
 - Ignorance of vertical mixing
 - Overemphasis of temperature effects

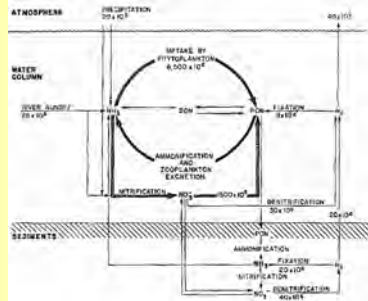
EEOS630

Slide 72 Brandt's strengths & weaknesses

NOTES:

The Nitrogen cycle

From Mills (1989) Fig. 23 (p. 56)



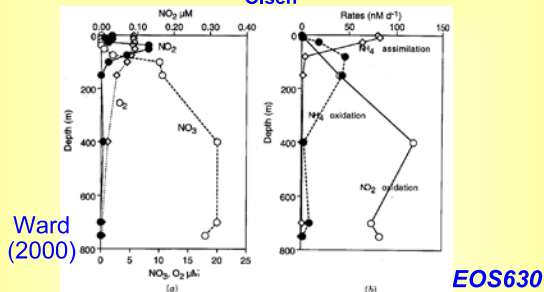
Fluxes
in
metric
tons per
year
EEOS630

Slide 73 The Nitrogen cycle

NOTES:

Subsurface NO₂⁻ maximum

Produced by light inhibition of NO₂⁻ oxidizers; Ward & Olsen

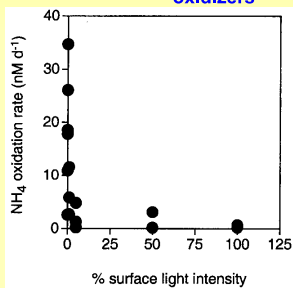


Slide 74 Subsurface NO₂⁻ maximum

NOTES:

Light inhibition of NH₄⁺ oxidation

Ward (2000), NO₂⁻ oxidizers more inhibited than NH₄⁺ oxidizers



Slide 75 Light inhibition of NH₄⁺ oxidation

NOTES:

Denitrification in the coastal zone

Seitzinger (1988): about 40-50% of N loading to coastal zones lost as N_2 .

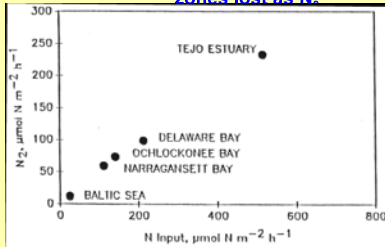


Fig. 2. Denitrification rates vs. external N inputs rates in estuaries. Data explained in Table 7.

EEOS630

Slide 79 Denitrification in the coastal zone

NOTES:

About 40-50% of N load lost as N_2

Seitzinger (1988); Giblin & Howes: similar rates for BH

Table 7. The importance of denitrification as a sink for external nitrogen inputs to various aquatic systems. Only systems in which denitrification rates were measured at near-ambient conditions are included. N Input, $\mu\text{mol N m}^{-2} \text{h}^{-1}$.

Location	N Input	N Input removed by denitrification (%)	N Input*	Time Interval
Lakes				
ELA 227f	3,200	1.4	1	Annual
Kvidet søf	1,227	7	2	Annual
Byrrup Langel	690	28	2	Annual
Arresø	81	14	3	Annual
Okeechobee	46	9-23	4	Annual
Meadowf	45	36	5	Annual
Rivers				
Potomac	632	35	6	Fall
Delaware	1,452	20	7	Summer
Estuaries				
Tejo estuary	316	43	8	Fall
Delaware Bay	213	46	9	Jul-Sep
Ochlockonee Bay	174	43	10	Annual
Narragansett Bay	112	50	11	Annual
Baltic Sea	25	40	12	Annual
Baltic Sea	25	25	13	Annual
Freshwater Bayf	25	20	14	Annual

0

Slide 80 About 40-50% of N load lost as N_2

NOTES:

Denitrification in Boston Harbor sediments: *Ampelisca* mats

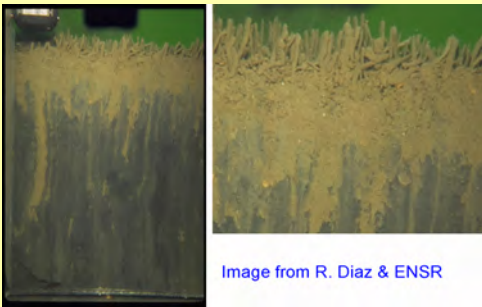


Image from R. Diaz & ENSR

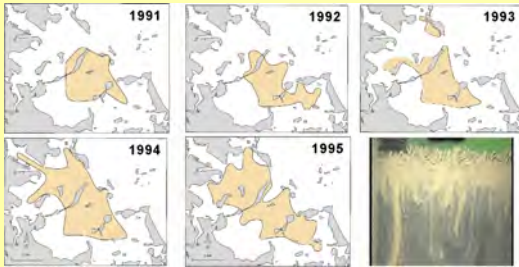
S630

Slide 81 Denitrification in Boston Harbor sediments: *Ampelisca* mats

NOTES:

Ampelisca mats in Boston Harbor

Oligochaete-spionid-Capitella → *Ampelisca*



Data from MWRA & ENSR

EEOS630

Slide 82 Ampelisca mats in Boston Harbor

NOTES:

Four major revolutions

In our understanding of nutrient limitation

- Brandt (1899) was correct to focus on N limitation, Liebig's law, and the role of denitrification, but he missed the role of vertical mixing providing vertical flux of nutrients
 - The anammox pathway, missed until 2003 provides further insight into the central role of nitrogen removal
- Chemostat work by Droop (1968), Caperon & Meyer (1972), Fuhs & Rhee revealed the central importance of the **Internal nutrient pool** in controlling μ
- Goldman (Goldman *et al.* 1979, 1980) argued that phytoplankton in nature tend to grow at high relative growth rates, otherwise they would not exhibit Redfield stoichiometry. The internal nutrient pool tends to follow Redfield stoichiometry.
 - Nutrient input controls phytoplankton biomass & species composition
 - One phytoplankton assemblage rapidly replaced by another, each with high relative growth rate.
- Martin's Iron hypothesis: iron is the Liebigian nutrient in major areas of the world's ocean

Slide 83 Four major revolutions

NOTES: